

National Centre for Sustainable Subsurface Utilization of the Norwegian Continental Shelf

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## Geochemical challenges of hydrogen storage in Salt caverns

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### Importance of H2 Economy

**O** Efforts to achieve energy transition and climate neutrality in EU and the rest of the world

O The European Hydrogen Backbone (EHB) initiative (2021):

- H2 Demand for transportation in 2050: 285 TWh ~ 12% of the total demand
  - 68 TWh for the Aviation sector
  - 217 TWh for heavy-load transport
- H2 Demand for electricity generation: 626 TWh in 2050 ~ 7% of the EU electricity demand

**O The International Energy Agency (IEA,2022):** 520 MT of clean hydrogen needed annually by 2050 to reach net zero



### H2 Storage importance

- TWh storage of H2: Only underground storage can provide enough capacity
- Other options have limited capacity



Fangxuan Chen et al., 2022



### Background

- O Advantages of salt caverns vs porous rocks for H2 storage
- Ductility of salt rock
- Impermeable: less prone to H2 leakage
- Less surface contact with rock minerals
- Less cushion gas requirement
- Less storage costs





#### Salt Formations and H2 Storage Capacity in Europe





#### **Geochemical Challenges**



Halite with sylvite (left core, NPD, 2022) and halite with discontinuous stringers of anhydrite and claystone (right core).

- Salt formations are usually associated with high amounts of impurities
- Impurities such as anhydrite, clays, carbonates, and iron that could react with hydrogen and generate unwanted gases
- The presence of Halophilic bacteria in the cavern could lead to the formation of hydrogen sulfide in the presence of dissolved sulfate resulting from the leaching.

# Challenges

- O Hydrogen is an electron donor for certain types of bacteria
- O The presence of sulfate-reducing bacteria in salt caverns has been proven (Laura Schwab et al, 2022)
- Anhydrite and gypsum dissociation during the leaching can release sulfate, which could react with H2 at the interface
- The catalytic reactions result in H2S production
- The presence of H2S in the gas phase is detrimental to the safety of the storage operation





# Chemical model in PHREEQC



- H2 is stored in the cavern at 180 bar
- H2 is diffused from the gas cap to the brine phase
- Sulfate is allowed to diffuse from the sump to the brine

Reaction: SO4-2 + 9H+ + 8e- = HS- + 4H2O

- The reaction kinetics is determined by microbial activity
- The kinetic rate depends on the environment
- In this model, k = 9e-8 mol/kgw/s (Herrera, L. et al. 1997)
- The Monod Equation for bacterial sulfate reduction:

$$H2S \ rate = k * \frac{[SO_4^{2-}]}{1E - 04 + [SO_4^{2-}]} * \frac{[H_2]}{1E - 04 + [H2]}$$
(Laban, 2020)



#### Cavern data (brine and sump)

Data are taken from C. Hemme, W. van Berk, 2017

Brine		Sump
Al 3.706e-	07	Al 3.706e-07
Ba 9.097e-	07	Ba 8.136e-07
C 7.077e-	03	C 2.901e-05
Ca 6.333e-	02	Ca 5.488e-02
Cl 6.310e+	00	Cl 6.306e+00
Fe 1.415e-	03	Fe 2.257e-03
K 1.010e-	04	K 1.010e-04
Mg 1.315e-	03	Mg 1.464e-02
Mn 9.100e-	07	Mn 9.100e-07
N 3.008e-	04	N 1.152e-03
Na 6.310e+	00	Na 6.306e+00
P 4.840e-	07	P 4.840e-07
S 6.262e-	02	S 7.108e-02
Si 3.368e-	05	Si 5.242e-05

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• pH of the brine: 5.7

# Equilibrium reactions reactions

Equilibrium phase	Equilibrium reactions	logK
Halite	NaCl = Cl- + Na+	1.570
Anhydrite	CaSO4 = Ca2+ + SO4 2-	-4.39
Siderite	FeCO3 = Fe2+ + CO3 2-	-10.89
Quartz	SiO2 + 2H2O = H4SiO4	-3.98
Barite	BaSO4 = Ba 2+ + SO4 2-	-9.97
Pyrite	FeS2 + 2H+ 2e- = Fe 2+ + 2HS-	-18.479
Dolomite	CaMg(CO3)2 = Ca 2+ + Mg 2+ + 2CO3 2-	-17.09
Mackinawite	FeS + H+ = Fe 2+ + HS-	-4.648
Sulfur	S + 2H+ + 2e- = H2S	4.882
Calcite	CaCO3 = CO3 2- + Ca 2+	-8.48

C. Hemme, W. van Berk, 2017



# Hydrogen Diffusion: Fickian diffusion

- O Hydrogen solubility calculated in Phreeqc at 50 C, 180 bar.
- **O** Initial amount of hydrogen in the brine = 1.465e-02 mol/kgw
- O Hydrogen diffusion in the brine and sump over time and space is defined using Fick's 2<sup>nd</sup> law of diffusion

**O** Fick's equation 
$$\frac{\partial c}{\partial t} = D\left(\frac{\partial^2 C}{\partial x^2}\right)$$

• Solution for constant surface concentration:  $C(x, t) = C_{sat} \operatorname{erfc}\left(\frac{x}{2\sqrt{Dt}}\right)$ 

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O Diffusion coefficient : 5.13e-9 m2/s (Hemme et al., 2018)

# Base case simulation results – 5 years scenario



- H2 concentration is constantly declining meaning that consumption gradually increases
- Sulfate concentration steadily declines during the first 25 months
- H2S generation is maintained thanks to sulfate diffusion to the brine from the 25th months
- Sulfate diffusion from the sump is therefore the main driver of H2S generation in the brine



## Effect on pH evolution





## Sensitivity



#### **Kinetic Rate effect**



pH effect

# FeS precipitation at high temperature

#### Sulfate reduction

 $4H2 + SO2-4 \rightarrow H2S + 4H2O$   $\Delta G_r^o = -20$  kJ mol−1 (Tori M. Hoehler et al., 2001)

#### Pyrite formation

Fe+2 + 2 HS- = FeS2 + 2H2  $\Delta G_r^o$  = -41 kJ mol-1 (Joana Thiel, 2019)



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## Pressure effect







#### Parametric Variation of H2S formation



# H2S mitigation methods

• Adding NaOH leads to pH increase, hence a reduction in H2S amount generated in comparison to the base case



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# Conclusion

- O Under optimum conditions, the presence of sulfate in the brine can lead to sulfate reduction and H2S production
- O Main Parameters influencing H2S generation: kinetic rate
- O H2S generation can be mitigated by increasing the pH through NaOH addition



#### Future Work

**O** Model Calibration

O Application to the Salt deposition in Norway



#### THANKS FOR LISTENING





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